



Research article

Mechanisms and age estimates of continental-scale endorheic to exorheic drainage transition: Douro River, Western Iberia



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ABSTRACT

In western Iberia, mechanisms that can explain the transition from endorheic to exorheic continental-scale drainage reorganization are foreland basin overspill, headwards erosion and capture by an Atlantic river, or a combination of both. To explore these, we have investigated the Portuguese sector of the Douro River, the locus of drainage reorganization. The Douro River is routed downstream through the weak sedimentary infill of the Douro Cenozoic Basin, after which the river cuts down through harder granitic and metamorphic rocks crossed by active fault zones, before reaching the Atlantic coast. We investigated the drainage reorganization using an integrated approach that combined remote sensing, field survey and geochronology, applied to Pliocene–Quaternary fluvial sediments and landforms. The older drainage record is documented by a series of high and intermediate landform levels comprising: (1) a high level (1000–500 m a.s.l.) faulted regional fluvial erosion surface, the North Iberian Meseta planation surface and the Mountains and Plateaus of Northern Portugal, recording the endorheic drainage of the Douro Cenozoic Basin; (2) a first inset level at 650–600 m a.s.l., comprising a broad fluvial surface developed onto a large ENE–WSW depression, interpreted as recording the initiation of the continental scale reorganization; and (3) an inset fluvial surface at 550–400 m a.s.l., corresponding to the establishment of the exorheic ancestral Douro valley. The younger drainage record comprises an entrenched fluvial strath terrace sequence of up to 9 levels (T9 = oldest), positioned at 246–242 m above the modern river base; T1 = youngest, positioned at +17–13 m. Levels T1 and T3 display localized fault offsets. The three lowest terrace levels (T3–T1) were dated using optically stimulated luminescence techniques with results ranging from > 230–360 ka (T3), through 57 ka (T2) to 39–12 ka (T1). Fluvial incision rates of the younger terraces were quantified and temporally extrapolated to model the ages of the intermediate to high elevation levels of the early drainage record. Integration of incision data informs on the probable timing of the drainage reorganization and the initial adjustment, ~3.7–1.8 Ma. This was followed by acceleration of incision, producing the entrenched river terrace sequence developed via spatial and temporal variations in rock strength, uplift and cyclic cool-climate variability as the river adjusted to the Atlantic base level.

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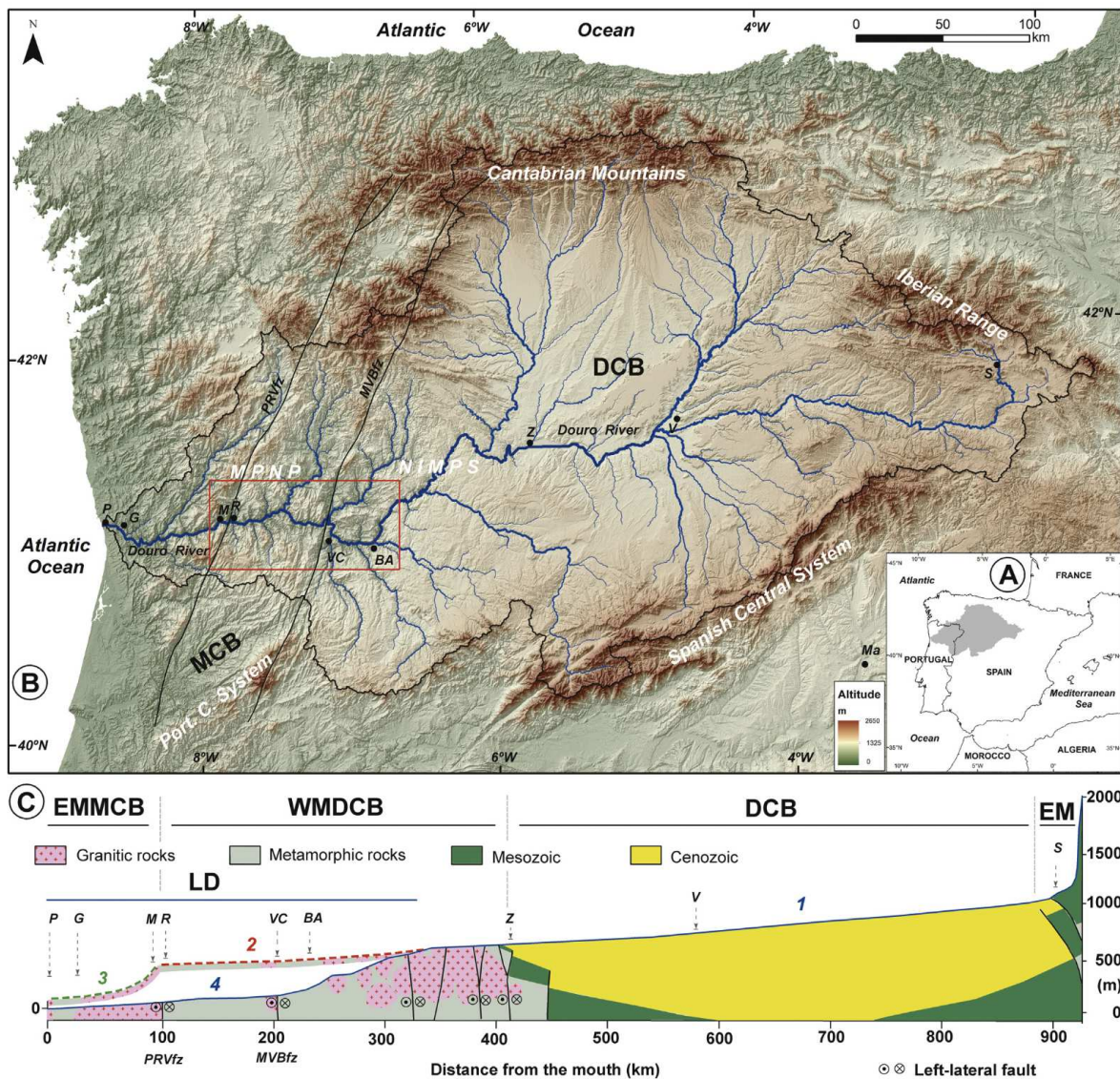


Fig. 1. The Douro catchment (drainage divide = black line) and its drainage network (blue lines) (B), inset showing location in Iberia (A) and longitudinal profile (C) with bedrock geology relationships. EM – eastern margin of the Douro Cenozoic Basin; DCB - Douro Cenozoic Basin; WMDCB - western margin of the Douro Cenozoic Basin; EMMCB - eastern margin of the Mondego Cenozoic Basin; MCB - Mondego Cenozoic Basin; MPNP - Mountains and Plateaus of Northern Portugal; NIMPS - North Iberian Meseta planation surface; MVBfz - Manteigas-Vilarica-Braganca fault zone; PRVfz - Penacova-Régua-Verín fault zone; LD - Lower Douro; 1 - long profile of the Douro River in the DCB; 2 - long profile of the ancestral Douro in the study area; 3 - long profile of an ancestral coastal Atlantic river; 4 - long profile of the modern Douro downstream of the DCB; P - Porto; G - Gondomar; M - Mesão Frio; R - Régua; VC - Vila Nova de Foz Côa; BA - Barca d'Alva; Z - Zamora; V - Valladolid; S - Soria; Ma - Madrid; Red rectangle on (B) - study area. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

1. Introduction

Continental areas with internal drainage adjacent to oceanic regions can reconnect with marine base level by (i) overspill of the internal drainage into an adjacent river system with marine base level connection (e.g., [Heidarzadeh et al., 2017](#)) or (ii) capture from adjacent river systems already linked to a marine base level (e.g., [Stokes and Mather, 2003](#)). Such switches from internal (endorheic) to external (exorheic) drainage can result in significant drainage network re-organisation,

including marked changes in drainage routing and substantial fluvial incision and erosion in the area where overspill or capture has occurred. This incision then propagates upstream into the former endorheic system as it equilibrates with the newly connected and lower marine base level.

The study of how rivers that cut transversely through mountains or plateaus has seen a significant resurgence in the last two decades, using field and modelling approaches (e.g., [Young and Spamer, 2001](#); [Stokes and Mather, 2003](#); [Twidale, 2004](#); [Douglass et al., 2009](#); [Larson et al.,](#)

2017). These studies show that the development of transverse drainage can often lead to re-organisation of catchments, allowing formerly internally drained basins to switch to externally draining systems. The process by which this transformation occurs is commonly attributed to the mechanisms of antecedence, superimposition, overflow (overspill) or piracy (capture). However, the timing and mechanisms, as well as the role of the possible interplay between surface and tectonic processes, are a matter of debate in most cases. This is because the drainage development process is typically dominated by erosion, resulting in an incomplete or missing record. As stated by Heidarzadeh et al. (2017), the typically proposed mechanisms are basin overspill (a top-down process) and headward erosion (a bottom-up process), but the lack of unequivocal diagnostic criteria for distinguishing between such mechanisms can hamper understanding. However, using integrated remote sensing, fieldwork, geochronological and modelling approaches in the study of fragmentary fluvial landscape elements can yield valuable insights into the mechanisms and rates at which transverse drainage developed.

Continental Cenozoic basins that underwent an endorheic to exorheic transition are reported in several Iberian river systems, encompassing a wide range of spatial and temporal scales (e.g., Mather, 2000; Stokes et al., 2002, 2018; Stokes, 2008; Soria-Jáuregui et al., 2018; Struth et al., 2019). These studies illustrate that the best interpretative model of the mechanism driving the endorheic–exorheic drainage reorganization should be based on the integration of detailed geological and geomorphological studies using the trunk river and, if necessary, its tributaries. Once the exorheism has been established, the patterns and rates of adjustment can be investigated using dated river terraces (e.g., Silva et al., 2017).

In this paper we focus on continental-scale drainage reorganization in an intraplate setting using the largest Iberian drainage system, the Douro River (Duero in Spain). The catchment of the Douro comprises a continental foreland Cenozoic basin that underwent an endorheic to exorheic transition. The complexity of the interpretation concerning the initiation and evolution of the Douro River was first presented by Lautensach (1932, 1987). Transverse drainage has been proposed to have developed from antecedence (erosional downcutting by an already existing river system) (Ferreira, 1978) or by drainage reorganization through headward erosion and capture (e.g., Feio, 1951; Ferreira, 1986; Martín Serrano, 1991; Antón et al., 2018; Struth et al., 2019). Overspill has not been explored hitherto as an explanation of Douro drainage evolution.

Most interpretations of the transition of the previously endorheic Cenozoic basins of central Iberia (e.g., Douro and Madrid Cenozoic basins) to Atlantic drainage systems have suggested that headward erosion by Atlantic fluvial systems captured the central Iberia drainage and eventually triggered drainage reversal (e.g., Gutierrez-Elorza and Pérez-González, 1993; Alonso-Gavilán et al., 2004 and references therein). However, other studies of Iberian river systems have shown that similar drainage evolution may be instigated by a combination of aqueous overflow of the basin and by capture-related headward erosion (e.g., Ebro Basin; García-Castellanos et al., 2003). Also, the transition from endorheic to exorheic drainage in the Madrid Cenozoic Basin is considered to have occurred by overspill into the Lower Tejo Cenozoic Basin, at ~3.7 Ma, as evidenced by a culminant sedimentary unit (Cunha, 1992a, 2019); the later stage of fluvial incision associated with the River Tejo (Tagus) began as recently as ~2 Ma (Cunha et al., 2016, Cunha et al., 2012; Silva et al., 2017).

The study area (Freixo de Espada à Cinta – Régua), comprising the upstream part of the Lower Douro catchment (Figs. 1, 2 and 3), is crucial for understanding the regional transition from endorheic to exorheic drainage, due to its location between the previously endorheic Douro Cenozoic Basin and the Atlantic drainage of the Mondego Cenozoic Basin (Cunha, 2019).

More broadly, the study area also allows the formation and destruction of a low-relief plateau morphology to be addressed. Changes

in climate or in the locus of tectonic activity can lead to disequilibrium conditions, a potential trigger for rapid destruction of the plateau morphology via fluvial incision (Meek and Douglass, 2001; House et al., 2008; Craddock et al., 2010; Heidarzadeh et al., 2017).

In several reaches of the study area, the Douro shows two high fluvial erosion levels at 650–600 m and 500–450 m above sea level (a.s.l.) that are inset into the regional planation surface and above a staircase of strath terraces (fluvially eroded benches that are etched into bedrock) (Ferreira, 1971, 1978). The lower-level strath terraces are capped by a veneer of fluvial sediment that enabled a degree of chronological control for the endorheic–exorheic transition process using luminescence dating. The geomorphological data and dating of the terrace deposits provides temporal estimates for fluvial aggradation and quantification of fluvial incision rates. The collective integration of terrace incision rates is then used to evaluate differential crustal uplift. Collectively, this work informs on the timing and the controlling mechanisms of the transition from endorheic to exorheic drainage in the Douro Basin, which were hitherto poorly constrained because of a lack of geomorphological studies integrating detailed fieldwork and absolute dating of sedimentary deposits and landforms.

2. Geomorphological and geological background

The Douro is one of the longest Iberian rivers (~927 km), with a source at 1700 m a.s.l. in the “Sierra de Urbión” (Iberian Range, Spain) and its mouth at Porto, on the Atlantic coast of Portugal (Fig. 1). The drainage basin covers approximately 97,290 km², with 80% in Spain and the remaining 20% in Portugal (e.g., Tockner et al., 2009).

The long profile gradient of the modern Douro River varies as it crosses different geological regions. In the Iberian Ranges, within its uppermost reach, the profile has a steep gradient coincident with Mesozoic bedrock (Fig. 1). Downstream, it has a classic ‘graded-profile’ shape as the Douro crosses through the sedimentary infill of the Douro Cenozoic Basin. Here, the river is incised by 20–200 m below the basin surface (e.g. Antón et al., 2012 and references therein). The trunk river and its tributaries display low-gradient, concave profiles with valleys characterized by well-developed river terrace staircases (e.g. Silva et al., 2017). The downstream sector along the western margin of the Douro Cenozoic Basin is an extensive low-to-high relief faulted area at 500–1000 m a.s.l., corresponding to the North Iberian Meseta planation surface (NIMPS) (Fischer, 1894; Hernandez-Pacheco, 1911; Dantin, 1912; Lautensach, 1932; Solé, 1952) and the Mountains and Plateaus of Northern Portugal (MPNP) (Ferreira, 1978, 1980; Galve et al., 2019) (Fig. 1). Downstream to the west of the Montemuro–Marão–Alvão relief barrier, the regional elevation is lower and records the ancestral drainage route towards the Mondego Cenozoic Basin (Figs. 1, 2 and 3).

Along the North Iberian Meseta planation surface, the Douro is incised into granite bedrock, forming a 150 km long knick zone. Here, it flows through deeply incised gorges (Arribas do Douro), with up to 600 m of incision recorded at the Águeda–Douro confluence (Antón et al., 2012, 2014, 2015, 2018). Along the Lower Douro (~230 km), the river flows approximately E–W from the Águeda confluence to the Atlantic coast. Its valley is incised into metamorphic and granitic basement, displaying an almost graded lower gradient, v-shaped form (Figs. 1, 2 and 3).

The Arribas do Douro knick zone (Figs. 1, 2 and 3) represents a resistant lithological threshold limiting upstream headward erosion from the Atlantic Ocean, thus acting as a local lithological base level for the upstream area, controlling the depth of incision and rate of headward erosion into the weaker sedimentary rocks of the Douro Cenozoic Basin (e.g. Martín Serrano, 1991, 1994).

The sedimentary infill of the Douro Cenozoic Basin mainly comprises Paleogene to Late Miocene fluvial to lacustrine arkoses, silty-clays and limestones the top unit of this succession is the Páramo Formation, dated as 9.7–9.6 Ma by Krijgsman et al. (1996), but also upper Tortonian to early Pleistocene alluvial fan gravels and sands (e.g.

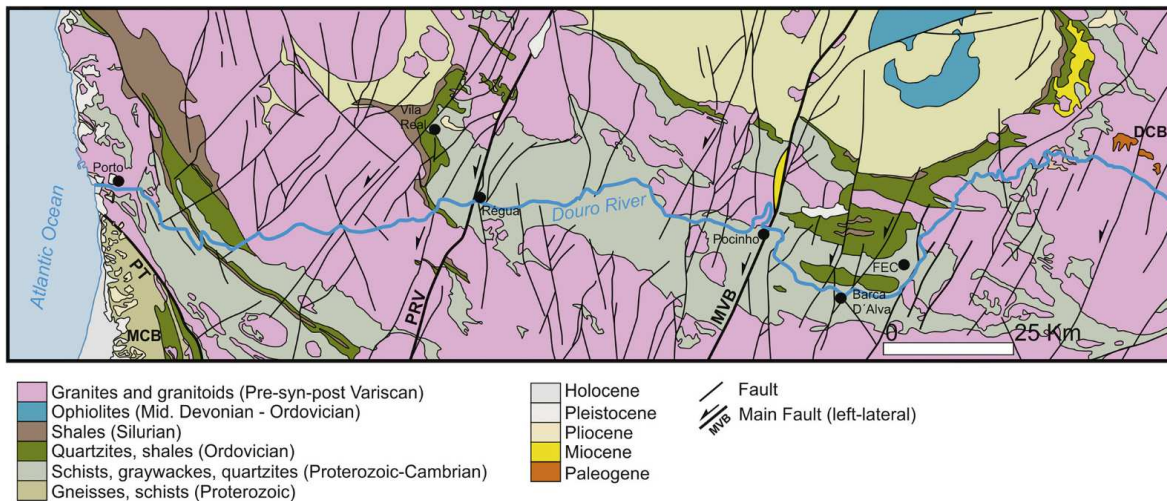


Fig. 2. The geology of the region crossed by the Douro River between the westernmost part of the Douro Cenozoic Basin (DCB) and the Mondego Cenozoic Basin (MCB) (adapted from IGME 1:1,000,000). The study area (Freixo de Espada à Cinta – Régua) is centred on the large linear outcrop of less resistant metamorphic rocks (mainly schists and greywackes; Proterozoic–Cambrian), located between resistant granites. FEC - Freixo de Espada à Cinta; MVBfz - Manteigas-Vilariça-Bragança fault zone; PRVfz - Penacova-Régua-Verín fault zone; PT - Porto-Tomar fault zone.

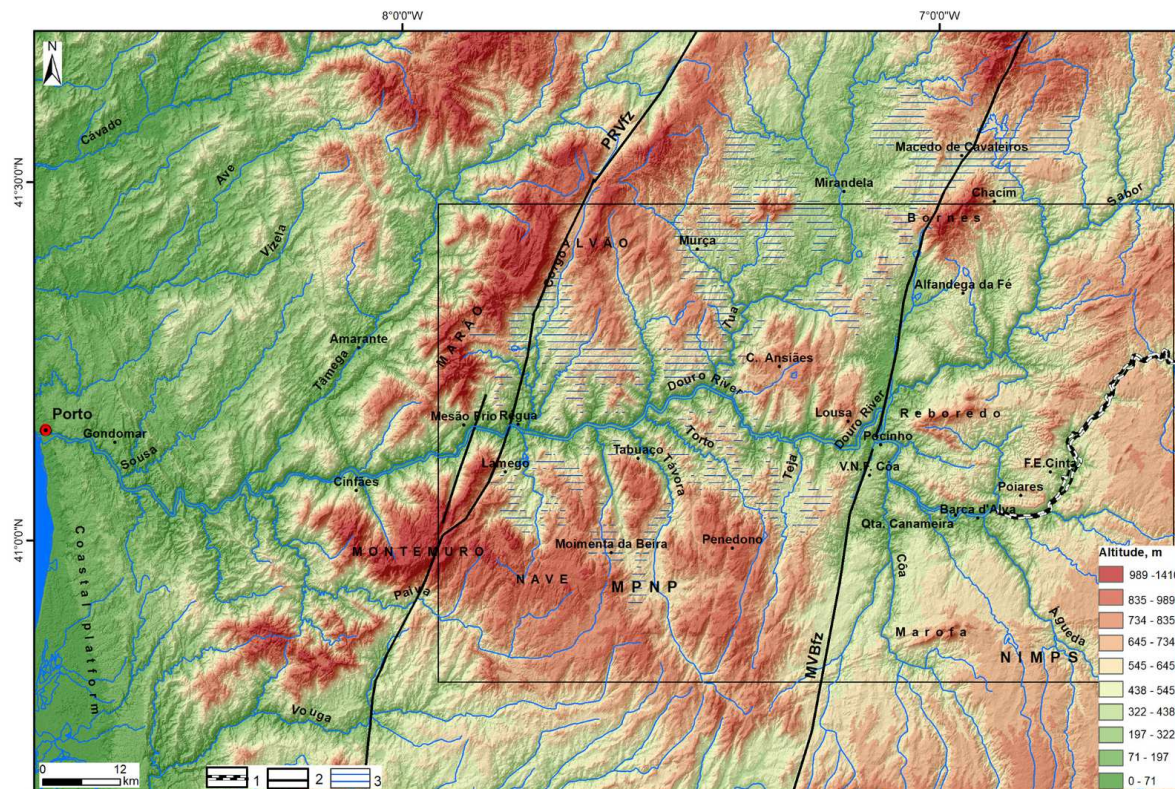


Fig. 3. Digital Elevation Model (DEM) showing, from east to west, the areas crossed by the Douro from the Spanish border to its mouth at the Atlantic coast. 1 - The large knick zone where the Arribas do Douro gorge is located; 2 - main fault; 3 - regional fluvial level at ~650–600 m a.s.l.; NIMPS - North Iberian Meseta planation surface; MPNP - Mountains and Plateaus of Northern Portugal (comprising the Nave, Montemuro, Marão and Alvão elevated areas); MVBfz - Manteigas-Vilariça-Bragança fault zone; PRVfz - Penacova-Régua-Verín fault zone. The inset indicates the study area.

Alonso-Zarza et al., 2002; Alonso-Gavilán et al., 2004; de Vicente et al., 2018; Cunha et al., 2019). The topmost unit of the sedimentary infill consists of relatively thin, but extensive, alluvial-fan ochre deposits ("Rañas") connected to fluvial deposits that are considered to be Late Pliocene–Early Pleistocene in age, recording the establishment of exorheic drainage and pre-dating the period of fluvial incision related with terrace formation (e.g., Pérez-González and Gallardo, 1987; Martín Serrano, 1988, 1991, 1994; Gutierrez Elorza et al., 2002; Pereira et al., 2015). The Douro has a maximum of 14 terraces below the alluvial-fan ochre deposits (Rañas) within the Cenozoic Douro Basin, positioned between +141–144 m (highest and oldest) and +8–10 m (lowest and youngest) Pérez-González, 1982. Based on a height–age transfer function, Silva et al. (2017) obtained an age of 2.3 Ma for the general valley downcutting in central Spain for basins draining to the Atlantic. In western and central Iberia, paleoclimate was mainly hot and dry during the Paleogene–Zanclean, but was humid during the late Zanclean–Gelasian (e.g., Barrón et al., 2010; Pais et al., 2012).

The study area comprises the western margin of the Douro Cenozoic Basin. Here, remains of a thin sedimentary cover over the Variscan basement documents Paleogene–Zanclean alluvial/fluvial units that collectively record (1) drainage routed towards the east (to the endorheic Douro Cenozoic Basin) and (2) a culminant unit of the sedimentary infill (Late Pliocene to early Pleistocene), only locally preserved, that records the ancestral Douro River flowing towards the west (to the Atlantic coast) (Martín Serrano, 1991, 1994; Pereira et al., 2000). In the area around Régua, Ferreira (1971, 1978) identified two high fluvial erosion levels at 650–600 m and 500–450 m a.s.l. (named 'high Atlantic Douro' River levels) that are inset into the regional planation surface. The younger, inset high-level surface remnants correspond to a large fluvial valley, with localized occurrences of fluvial sedimentary deposits preserved within the Mirandela area (Fig. 3). Most of the downstream reaches of the Lower Douro show a form of the ancestral Douro valley, also with localized sedimentary deposits, such as at Porto (122 m a.s.l.; +120 m; Araújo et al., 2003) and Gondomar (136 m a.s.l.; +136 m) (Figs. 1 and 3). These ancestral relicts of Atlantic drainage are considered to span ~3.7–1.8 Ma (allostratigraphic unit UBS13; Cunha, 1992a,b, 2019), based mainly on palynology and stratigraphical/geomorphological correlation (Costa and Teixeira, 1957; Teixeira et al., 1962; Cachão, 1990; Pereira, 1997; Pais et al., 2012). A major climate change occurred at ca. 3.7 Ma, from the generally dry and hot climates prior to the Zanclean (Lower Pliocene) to the humid and hot climate during the latest Zanclean to Piacenzian. This major change is supported by studies of paleontology (Vieira et al., 2011, 2018; Silva et al., 2010) and sedimentology (Cunha, 1992a, 2000; Cunha et al., 1993; Pais et al., 2012) studies. The warm-humid climate during 3.7–2.8 Ma was followed by progressive cooling (Vieira et al., 2018). In summary, the ancestral Douro was identified in the study area to be routed towards the Atlantic Ocean, probably between ~3.7 and 2 Ma, when sea level reached a maximum of about +20–40 m (e.g. Miller et al., 2005), the climate was humid and hot, and the regional uplift rate was moderate (Cabral, 2012). Due to its culminant position within the fluvial landscape, where it is represented as a wide valley-floor surface at the top of the younger entrenched fluvial strath terrace sequence, this sedimentary unit/geomorphic reference-level thus pre-dates the period of enhanced fluvial incision.

The different stages of drainage evolution lie in an area crossed by NNE–SSW-oriented late Variscan fault zones. These structures were reactivated during the late Cenozoic as left-lateral strike-slip deformation zones (e.g., de Vicente et al., 2008, 2011, 2018; Cunha et al., 2019). In the Manteigas–Vilariça–Bragança fault zone, the small Vilariça and Longroiva tectonic basins were developed (Cabral, 1989; Cabral, 1995; Pereira and Azevedo, 1995). Although tectonic structures exist within the study area, the uplift rate during the last ~3 Ma is considered to have been low to moderate at most. As such, the landscape only shows localized adjustments in proximity to the major tectonic structures with little impact upon regional slope.

3. Materials and methods

The information presented here is derived from geomorphological, tectonic, stratigraphical, sedimentological and chronological data using a standard approach for fluvial landscape development studies (e.g. Stokes et al., 2012). Firstly, a geomorphological study of the broader region was undertaken to characterize the major fluvial surface units. This used an integrated approach of satellite remote sensing and Digital Elevation Model (DEM) analysis within GIS. The regional analysis was supplemented by more detailed remote sensing/DEM interpretation and identification of surfaces and terraces in areas critical to the understanding of drainage evolution. Secondly, river terrace levels were differentiated on the basis of (1) relative height, (2) lithostratigraphic and sedimentological characterization of the deposits and (3) luminescence dating of the terrace sediments. Collectively, these enabled quantification of the timing of fluvial aggradation and incision episodes.

Geomorphological mapping was undertaken in three stages: (1) analysis of satellite imagery of Landsat 7 and 8 (true colors), 1/25,000 black/white aerial photographs and two Digital Elevation Models, one based upon a 1/25,000 topographic database and another based on the ALOS data (<https://www.eorc.jaxa.jp/ALOS/en/aw3d30/index.htm>), informed by the DEM recommendations of Boulton and Stokes (2018); (2) production within GIS of geomorphological and fault-pattern maps; (3) field ground-truthing, using topographical (1/25,000) and geological (1/50,000) base maps.

Fieldwork included detailed geomorphological mapping, identification of probable active faults, stratigraphic logging and sedimentological characterization of fluvial deposits in order to obtain data on the depositional facies and the sedimentary processes and depositional environments they represent. This included sediment colour, texture, maximum particle size, clast lithology, fossil content, bedding and its overall depositional architecture.

For optically stimulated luminescence (OSL) dating, thirteen samples were collected from outcrops of river-terrace sedimentary units within the study area. The OSL dating technique, in terms of laboratory material preparation, measurement and age modelling, is described in full within the Supplementary Information. However, only the three lowest terraces, which record the more recent episodes of Douro River incision, had sand deposits suitable for sampling. The strategy involved targeting units from the base to the top of a given terrace level, in order to provide the timing of aggradation (e.g., Cunha et al., 2017). However, in some terrace deposits the target sands were rare or absent. Sampling tubes were hammered into previously cleaned outcrops. Immediately adjacent to each tube, a sub-sample of sediment was collected for determination of field water content, saturation water content and dose rate, all of which are required for the final age modelling.

4. Results

4.1. Old landform levels

In the study area (Fig. 4), several landform levels occur: (1) the North Iberian Meseta planation surface, (2) the Mountains and Plateaus of Northern Portugal, (3) some highly resistant quartzite uplands, (4) large depressions developed in less resistant metamorphic rocks, (5) remnants of low-resistance Cenozoic sedimentary cover, (6) high-level records of early Douro drainage, routed to the Atlantic, and (7) the entrenched lower part of the Douro valley, with river-terrace staircases.

The regional landscape is dominated by an extensive planation surface, with a slight regional slope towards the east. This surface is well preserved in granite basement but deeply dissected in less resistant metamorphic basement (Ferreira, 1971, 1978, 1980, 1991). The planation surface was generated during Cretaceous to late Miocene times, providing the sediments for the progressive infilling of the Douro Cenozoic Basin (to the east). Since the middle Tortonian (~9.6 Ma;

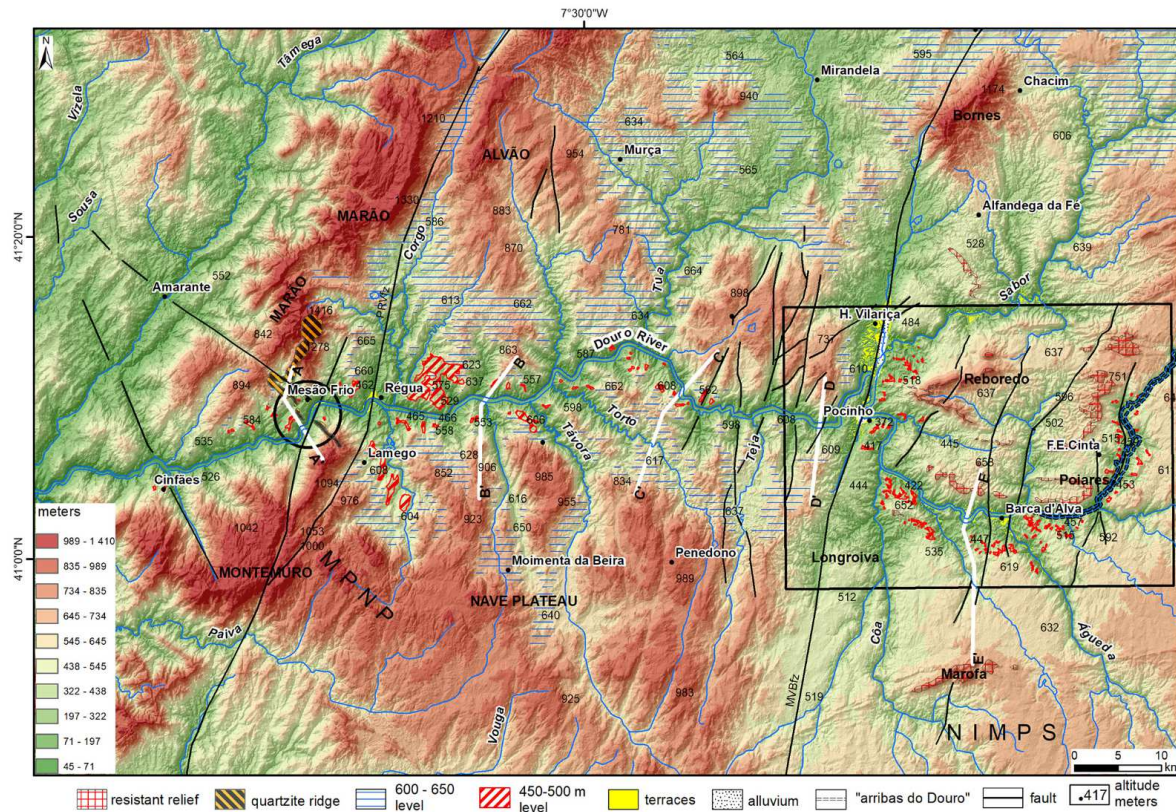


Fig. 4. DEM showing the hypsometry of the study area. It can be seen that the North Iberian Meseta planation surface (NIMPS) to the east of the MVBfz is tilted westwards (at 700 m to 500 m a.s.l.), but the planation surface rises westwards by tectonic steps up to 700–1000 m a.s.l. in the MPNP. A very wide fluvial level, at 600–650 m a.s.l., is represented. There are remnants of the large valley of the ancestral Douro at ~450–500 m a.s.l. The entrenched Douro valley exhibits incised meanders and reaches that were controlled by faults. The black circle indicates crossing (at Mesão Frio), by the Douro, of the Montemuro–Marão upland. The inset indicates the area shown in Fig. 6. A–A' to E–E' are topographic profiles of which are shown in Fig. 5.

Pereira, 1997; Cunha and Pereira, 2000; Cunha et al., 2000, 2019; Feio and Daveau, 2004; Pais et al., 2012), fault controlled differential uplift has segmented this planation surface. East of the Manteigas–Vilarica–Bragança fault zone (MVBfz), the planation surface is tilted from ~750 m a.s.l. (in the east; Miranda do Douro) to 500 m (in the west; at Vilarica and Longroiva tectonic depressions). In this sector the surface is referred to as the North Iberian Meseta planation surface (NIMPS). Between the Manteigas–Vilarica–Bragança and the Penacova–Régua–Vérin fault zones, the planation surface is located at between 657 m a.s.l. (in the east) and 700–1000 m a.s.l. (in the west), and is referred to as the Mountains and Plateaus of Northern Portugal (MPNP) (Figs. 3 and 4). The approximate NNE–SSW trend of the Montemuro and Marão mountain uplands formed a divide between the former easterly (towards the endorheic Douro Cenozoic Basin) and westerly (to the Atlantic Mondego Cenozoic Basin) routed drainage.

Resistant ridges, mainly comprising Ordovician quartzites, stand out above the regionally extensive plateaus developed on the less resistant metamorphic basement (Fig. 4). Examples are the Mount Reboredo (966 m a.s.l.), situated east of the Vilarica depression, the Poiars ridges (711 m a.s.l.), north of the Douro valley in the Barca d'Alva sector, and the Mount Marofa (976 m a.s.l.), at south of the Douro River and rising to ~280 m above the surrounding plateaus (NIMPS).

In the study area, several small remnants of Cenozoic sedimentary cover overlying the Variscan basement are preserved within and around the Vilarica and Longroiva tectonic depressions. During the late

Cenozoic, these small strike-slip tectonic basins were subjected to asymmetric subsidence, as evidenced by the eastward tilt of the Palaeogene to Miocene sequences (Cabral, 1989; Pereira and Azevedo, 1995; Cunha and Pereira, 2000) and by the preferential development of fluvial terraces on their western side.

The two highest Douro River levels, which are inset into the regional planation surface, were originally identified by Ferreira (1971, 1978) in the area around Régua at 650–600 m a.s.l. and 500–450 m. Our new mapping has extended the distribution of these highest Douro River levels over a much larger area (Figs. 3, 4 and 5). Both fluvial levels are discontinuous in the landscape due to erosion and display undulating surface expressions related to lithological contacts and/or tectonic displacement. In the study area, the first inset level occurs at 650–580 m a.s.l. and comprises a broad fluvial surface developed over a large ENE–WSW-oriented depression. The second high level (Fig. 1C – log profile 2) forms a discontinuous occurrence along the Douro valley margins, between 575 and 400 m a.s.l., being lower as 365 m a.s.l. at the Vilarica tectonic depression (Figs. 4 and 5). It corresponds to a wide fluvial valley-floor (up to ~5 km width). This surface does not occur in areas of higher topography west of the Manteigas–Vilarica–Bragança fault zone or where residual uplands occur. For example, it disappears where the high topography of the Carrazeda de Ansiães is traversed, downstream of the small Vilarica basin, and in the Cinfães–Lamego reach, where the Marão–Montemuro mountains are crossed (Figs. 3 and 4). East of the Manteigas–Vilarica–Bragança fault zone, this level is

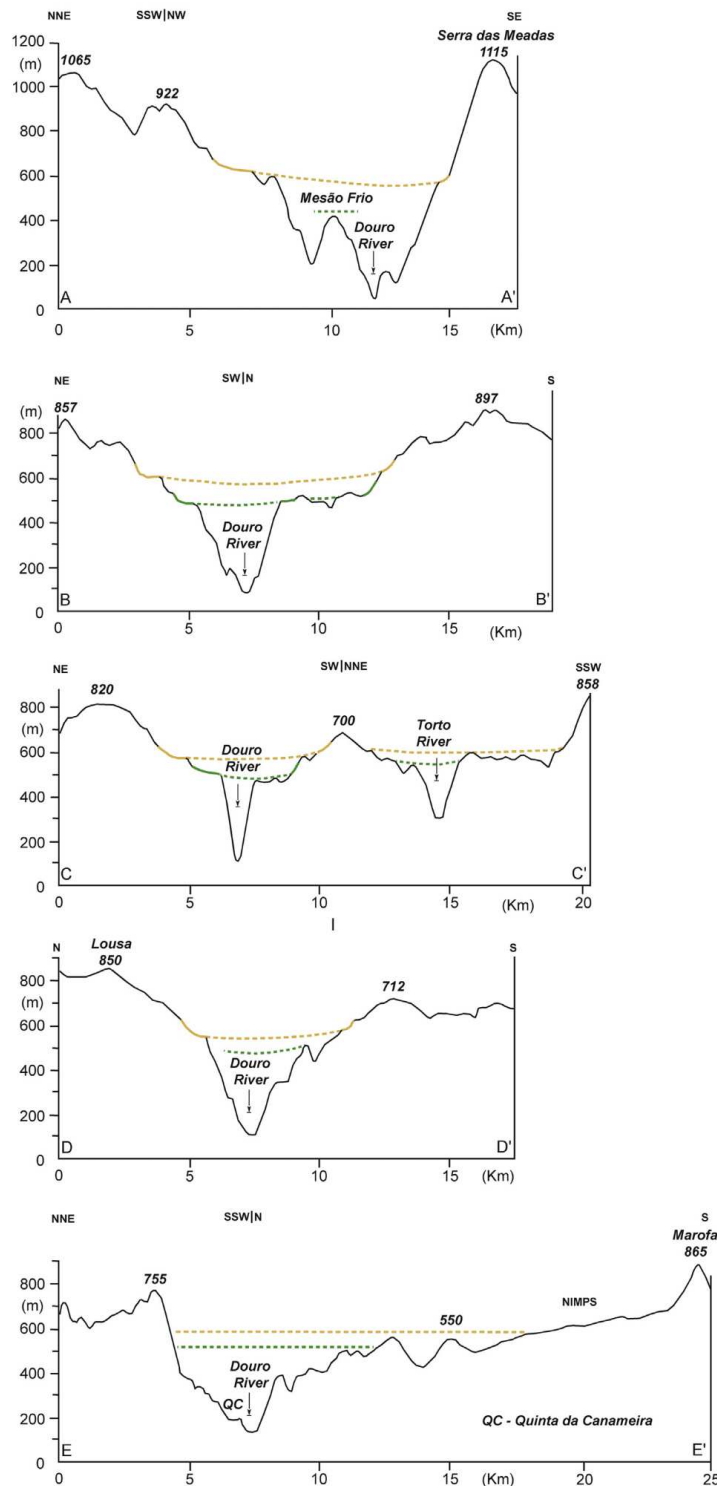


Fig. 5. Topographic profiles, transverse to the Douro valley, in the study area (for location see Fig. 4). The ~600 m a.s.l. level is represented by a dashed ochre line and the ~500 m a.s.l. level is represented by a dashed green line. At Mesão Frio, the total incision, from the 600 m a.s.l. level to the modern river bed, is 600 m. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

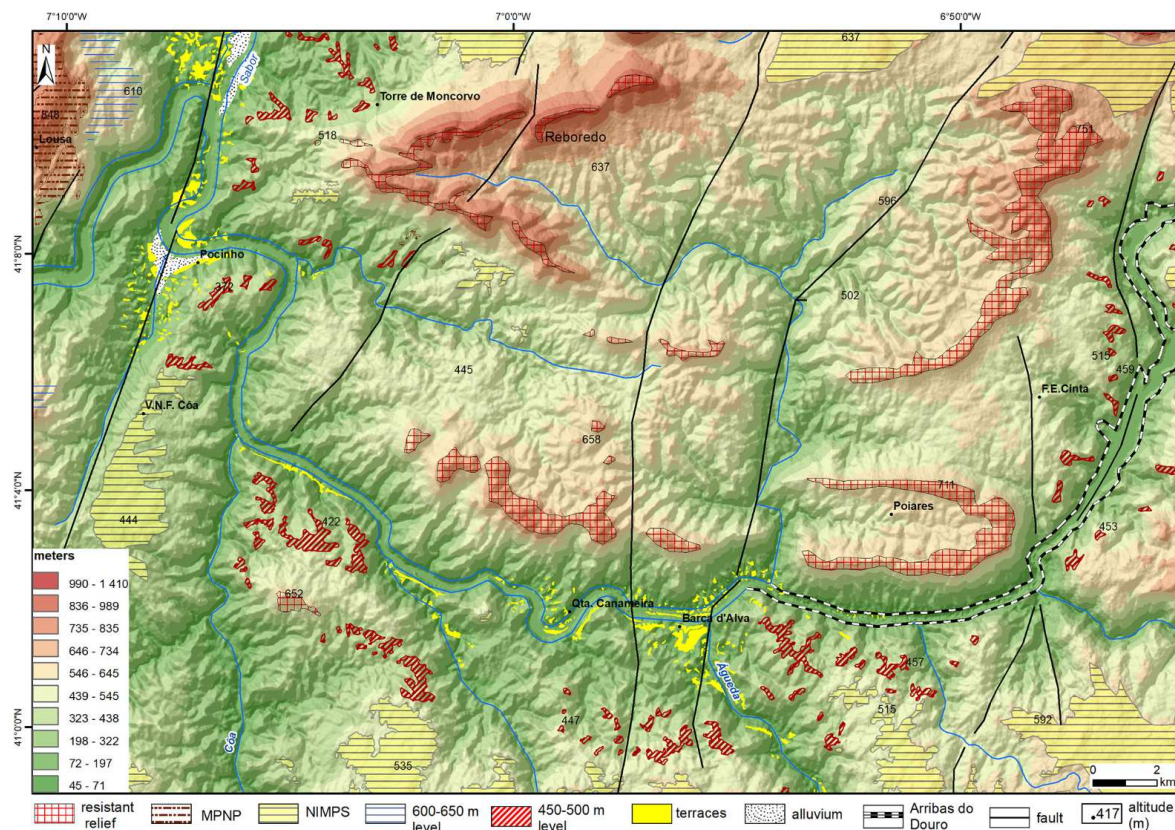


Fig. 6. Geomorphological map of the main terrace remains along the Douro valley, in the Freixo de Espada à Cinta – Vila Nova de Foz Côa area.

developed at a slightly lower altitude, namely 365 m a.s.l. at Pocinho, 455 m a.s.l. at Barca D'Alva and 515 m a.s.l. at Freixo de Espada à Cinta. This fluvial level is not represented further upstream until Zamora (Fig. 1). The narrow and entrenched parts of the valleys associated with incision by the Douro and its tributaries are located at lower levels than the remnants of this wide ancestral valley-floor.

4.2. River terrace staircases

4.2.1. Geomorphological characterization of the strath levels

In the Lower Douro, staircases of strath terraces are developed in sectors associated with (1) tectonic depressions, (2) less resistant lithologies (e.g. Cenozoic deposits) or (3) along the convex banks of enlarged meanders. The river has developed terraces of considerable extent within the Vilarica and Barca d'Alva reaches and on the inner convex side of the Quinta da Canameira valley meander (Figs. 6 and 7). Within the Quinta da Canameira meander, lateral migration caused fluvial sediments to be accumulated on the inside bank. Another location favourable for fluvial aggradation and terrace formation occurs near an unusual junction of the Sabor tributary with the Douro. Here, the Sabor flows in a direction opposite to the trunk river, promoting damming conditions near its mouth, locally known as Rebofa and especially intense during high floods. In addition, the marked narrowing of the Douro valley immediately downstream of the Sabor confluence, where it begins to incise into granite basement, creates natural damming conditions in this gorge reach, promoting aggradation and fill-terrace formation in the vicinity of the Sabor confluence. This confluence occurs within the elongated Vilarica tectonic depression,

further highlighting this sector as favourable for terrace formation.

In the reaches of the Douro valley, located between Barca d'Alva and Foz do Sabor (Fig. 6), up to nine strath terrace levels occur below the 455–365 m a.s.l. level. They are termed T1 to T9 (from the lower to the higher). The substantial lateral upstream–downstream extent of the T1 to the T6 straths means they are easily correlated. Terrace level assignment and correlation becomes challenging above the T6 level because preservation is poorer. Terrace T9 reaches up +246 m at Barca d'Alva and +220 m at Pocinho, within the Vilarica strike-slip basin (Table 1). The vertical spacing between the strath levels varies from ~15 to 25 m (average 20 m) for the levels below T7, whereas for the two higher T8 and T9 straths the vertical separations are larger (35–60 m). Strath levels below T6 are subparallel to the modern river bed, while T7, T8, T9 and the ancestral Douro valley-floor surface (the 455–365 m a.s.l. level) are steeper in the Quinta da Canameira and Barca d'Alva areas than their equivalents in the Pocinho area (Table 1; Figs. 8, 9, 10 and 11), i.e. showing a diverging upstream pattern.

The T1, T2 and T3 strath terraces have sediment thicknesses of up to 11 m. In contrast, T4 to T9 consist of erosional bedrock surfaces, some of them with only a thin cover of gravel (with boulders and cobbles). The best preserved and most visible terrace along the Douro valley is T3 (~+50 m a.r.b.). There is an equivalent of this terrace along the Sabor tributary, extending from the confluence with the Douro for > 30 km upstream (Martins et al., 2018). The extensive spatial development and sedimentary thickness of T3 contrasts with the lower and younger bedrock terraces, which have much less expression in the landscape.

The Quaternary deposits of the studied Lower Douro reach show different sedimentary characteristics (Table 2).



Fig. 7. A - Panoramic view, from the north, of the southern flank of the Douro valley at Barca D'Alva. The T3 terrace, on the left-hand side of the valley, can be seen. B - The Quinta da Canameira terrace staircase (Douro right bank) and the Poiars syncline (southern flank) quartzite ridges at the top. C - The Douro valley, Quinta do Vale Meão terrace staircase (left bank). D - Northwards view of the Douro valley near Pocinho; the Vilarica tectonic depression is the flat lowland in the background.

4.2.2. Luminescence ages of the strath terrace deposits

Only the three lower terraces have sufficient sediment thickness and suitable materials to be used for OSL dating. The OSL technique revealed ages of between > 360 – 230 ka and 13 ka (Tables 3 and 4). The samples from T3 have D_e values lying above $2x D_0$ (corresponding to 86% of luminescence saturation). In light of the possibility of small systematic errors during measurement of the luminescence signal, we adopted the approach of Wintle and Murray (2006) of only presenting minimum doses (equivalent to $2x D_0$), and thus minimum ages.

The T1 terrace was dated (Qz-OSL) using samples collected from a trench opened at Quinta do Vale Meão (Pocinho) that exposed silty sands: 102211 (25.2 ± 1.3 ka) was collected at ~ 1 m above the base of the T1 sediments, 102212 (17.0 ± 1.0 ka) at a middle position and 102219 (16.7 ± 1.0 ka) near the top. The T1 deposits were also dated from two samples collected from the topmost deposits at Cortes da

Veiga (Pocinho): 102220 (14.8 ± 1.0 ka) and 102218 (12.9 ± 0.8 ka). An aeolian brown coarse silty level covering the T3 terrace at Barca d'Alva (Douro left bank: $+ 65$ m a.r.b.) was dated 31 ± 2 ka (sample 102214). Previously, in another trench at the Vale Meão site, five fluvial sand samples provided OSL ages in the range 17 – 14 ka and the colluvium at the top was dated 14 – 12 ka, while an alluvial sequence with its surface at 120 m a.s.l. (T1) at Vilarica was dated 39 – 18 ka (Rockwell et al., 2009). In summary, the aggradation of T1 sediments represents the age range 39 – 12 ka.

Sample 102217, collected from a sand lense in the T2 gravel deposits at Ribeira da Barca (Pocinho), yielded an age of 57 ± 4 ka (Quartz-OSL). The age obtained by the pIRIR protocol (98 ± 7 ka) was considered to be an overestimate due to partial bleaching. Since the sample was collected from the basal deposits it is likely that the T2 aggradation in this area spanned ~ 60 – 50 ka.

Table 1

Elevations of the bedrock terrace levels (T1 to T9) in three studied reaches of the Douro valley. The fluvial 455–365 m a.s.l. level, corresponding to a large paleovalley, is here interpreted as the valley of the ancestral Douro River (ADR) (Atlantic drainage).

| Geomorphic references | Pocinho | | Qta. da Canameira | | Barca D'Alva | |
|----------------------------------|--------------|---------------------|-------------------|---------------------|--------------|---------------------|
| | Altitude (m) | Elevation a.r.b (m) | Altitude (m) | Elevation a.r.b (m) | Altitude (m) | Elevation a.r.b (m) |
| High and large paleovalley (ADR) | 365 | +262 | 450 | +334 | 455 | +337 |
| T9 | 323 | +220 | – | – | 364 | +246 |
| T8 | 289 | +186 | 319 | +203 | 320 | +202 |
| T7 | 229 | +126 | 273 | +157 | 291 | +173 |
| T6 | 209 | +106 | 227 | +111 | 240 | +122 |
| T5 | 195 | +92 | 207 | +91 | 210 | +92 |
| T4 | 179 | +76 | 186 | +70 | 190 | +72 |
| T3 | 151 | +48 | 168 | +52 | 171 | +53 |
| T2 | 134 | +32 | 150 | +34 | 145 | +27 |
| T1 | 116 | +13 | 132 | +16 | 135 | +17 |
| Modern river bed | 103 | 0 | 116 | 0 | 118 | 0 |

The T3 level was dated using the pIRIR protocol, since the Quartz-OSL signal was found to be saturated for all samples. Several minimum ages were obtained: 112233 (> 360 ka), from the basal T3 sediments at Barca d'Alva; 102215 (> 230 ka) and 102216 (> 290 ka), from the basal T3 deposits at Pocinho; and 102213 (> 270 ka), from the middle of the T3 depositional sequence at Barca d'Alva. Thus, the T3 sediment aggradation probably occurred between ~370 and 220 ka. It is interesting to note that Acheulian (Lower Paleolithic) artefacts were collected from the T3 terrace at Quinta do Vale Meão and also from a correlative Côa River terrace, located at +40 m a.r.b., at Penascosa (Zilhão, 1997). In central Portugal (River Tejo/Tagus), Acheulian industries associated with river-terrace deposits were recently dated to between ~360 and 200 ka (Cunha et al., 2017).

4.3. Tectonic activity

Field surveys allowed the identification of several faults affecting terrace deposits. At Quinta do Vale Meão, a fault (N20° E, 80° NW) marks the contact between granite, to the west, and T3 deposits (sandy matrix-supported gravels) to the east (Cabral, 1989, 1995). Given the vertical offset of ~20 m, and assuming a probable age of ~220 ka for the uppermost terrace deposits, an average slip rate for the vertical component of this fault can be estimated as ~0.09 m/ka. This location corresponds to the Vilariça fault, which delimits the eastern side of the minor Vilariça strike-slip tectonic basin (Figs. 9 and 10). Paleoseismic studies have shown T1 to have been displaced by the Vilariça fault (Rockwell et al., 2009; Cabral et al., 2010; Perea et al., 2010; Vilanova et al., 2014).

4.4. Fluvial incision and quantification of crustal uplift

The height of a fluvial terrace surface above the modern river profile (assuming negligible surface erosion) can be used to calculate the long-term fluvial incision rate, which in turn be used as a proxy for crustal uplift (Bridgland, 2000; Maddy et al., 2000), albeit taking into account the uncertainty that the river might not have fully achieved a steady-state longitudinal profile following the uplift.

Fluvial incision in the study area can be estimated using different geomorphic markers: (i) the North Iberian Meseta planation surface, (ii) the 455–365 m a.s.l. level (the ancestral Douro valley-floor) and (iii) the OSL-dated T1, T2 and T3 levels. Using the NIMPS as the oldest surface marker, we can estimate the differential uplift that has occurred since ~9.6 Ma (the beginning of relief differentiation; Pais et al., 2012). Faults have segmented this planation surface into several plateaus (Figs. 4, 6, 12 and 14). Using the long-term incision rate, immediately east of the Vilariça Fault (Pocinho - Barca d'Alva) the remnants are at 500–555 m a.s.l. (+367–+439 m a.r.b.), giving an incision rate of 0.04–0.05 m/ka. Immediately west of the Vilariça Fault, the corresponding remnants are at 607 m–845 m a.s.l. (+505–+743 m a.r.b.), giving an incision rate of 0.05–0.08 m/ka.

Using the ancestral Douro valley surface (the 575–365 m a.s.l. level) as a geomorphic reference (Figs. 4, 6, 9 and 11) allows estimation of the amount of fluvial incision since ~1.8 Ma (Fig. 13). Immediately east of the Vilariça Fault, within the Pocinho-Barca d'Alva sector, this relict valley surface is at ~364 m to 455 m a.s.l. (~+262–+337 m a.r.b.), providing an incision rate of ~0.15–0.19 m/ka. West of the Manteigas–Vilariça–Bragança fault zone, until Régua, the relict valley-floor surface is at 475–400 m a.s.l. (+400–+370 m a.r.b.), giving an incision rate of ~0.22–0.21 m/ka.

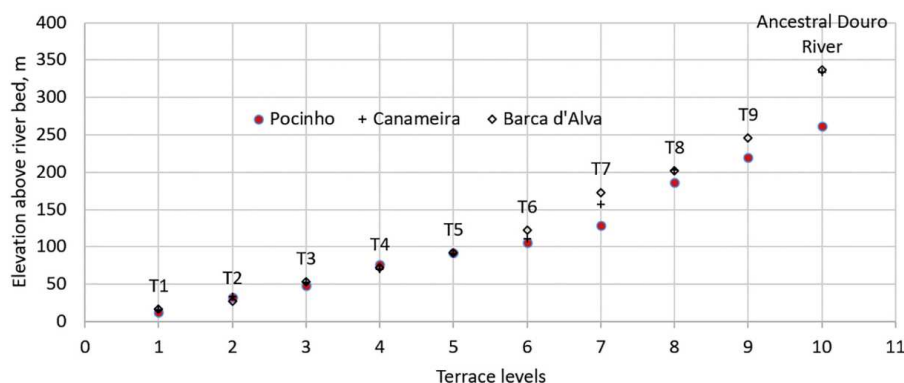


Fig. 8. Projection of the elevation (a.r.b.) of each terrace level, for three studied reaches of the Douro located between Barca d'Alva and Pocinho.

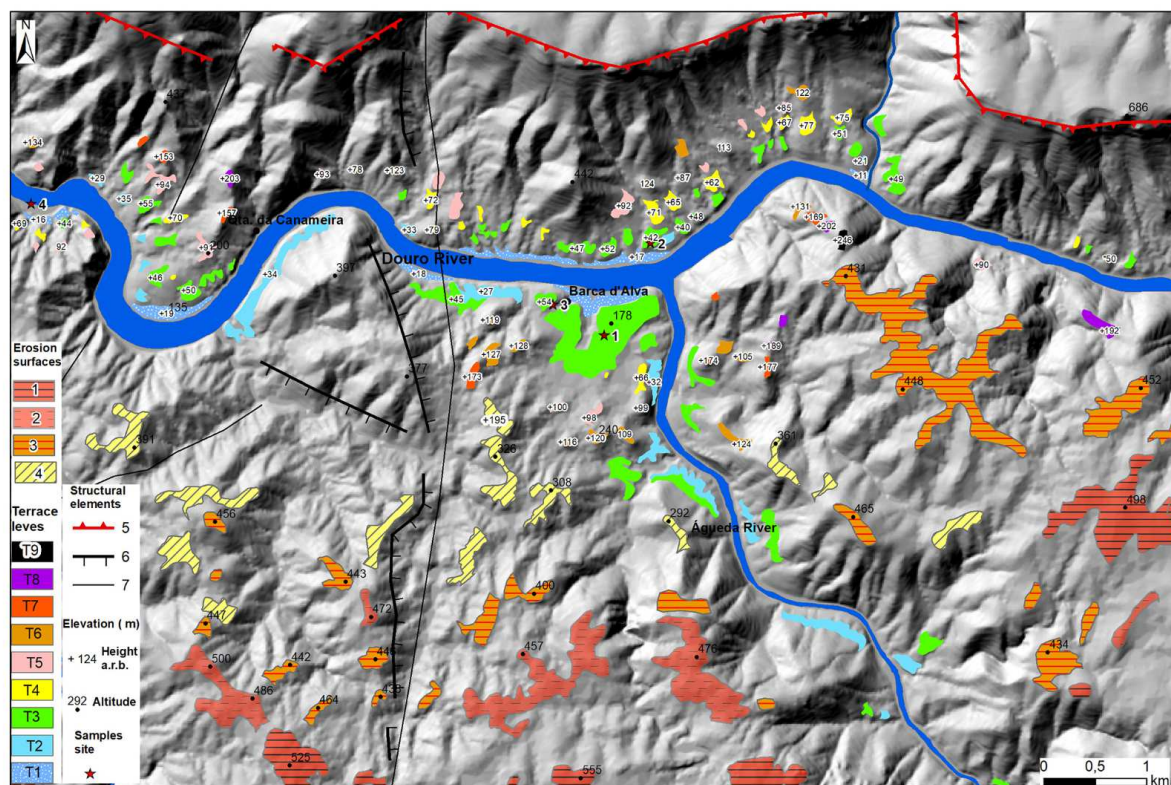


Fig. 9. Geomorphological map of the Barca d'Alva to Quinta da Canameira terrace staircases. 1 - North Iberian Meseta planation surface (NIMPS); 2 - degraded NIMPS; 3 and 4 - fluvial erosion surfaces at 430–465 m a.s.l. probable remnants of the ancestral Douro valley; 5 - scarp of structural relief; 6 - fault scarp; 7 - tectonic lineament; sampled sites (numbered stars): 1 - aeolian cover on T3, left bank (102214); 2 - T3, right bank (102213); 3 - T3, left bank (sample 112233); 4 - modern sample (102234).

The three OSL-dated lower terraces can also provide short-term incision rates, with reference to their terrace surface elevations and ages. Using the T3 level, the incision rate is -0.21 m/ka (48 m/230 ka) for Pocinho and -0.24 m/ka for Barca D'Alva. For T2, the incision rate is 0.44 m/ka (25 m/57 ka) for Pocinho and 0.54 m/ka for Barca d'Alva. Using T1, the incision rate is 1.40 m/ka (18 m/13 ka) for Pocinho and 1.41 m/ka for Barca D'Alva; this suggests a progressive increase during the recent Pleistocene.

By plotting the OSL ages derived from sediments forming the uppermost parts of T1–T3, and their elevations a.r.b., the probable ages of the older terraces T4–T9, and of the ancestral Douro valley-floor (ADR), can be extrapolated (Fig. 13).

5. Discussion

The transition from the erstwhile internal drainage of the Douro Cenozoic Basin (endorheic stage) to Atlantic Douro River drainage (exorheic stage) has been attributed primarily to antecedence or capture-related headward erosion (e.g., Ferreira, 1978, 1986; Martín Serrano, 1991; Antón et al., 2018; Struth et al., 2019). This transition occurred over a significant time span, encompassing the Pliocene–Pleistocene and reflecting interplay between spatially and temporally variable regional tectonic activity and climatic fluctuation, acting upon a landscape with marked variations in rock strength (e.g. basement versus basin-fill). We first explore these controls very broadly at a continental scale and highlight their potential roles for reorganization of the Douro system. Then we consider these controls more explicitly

and explore the evidence from different endorheic and exorheic sectors to interrogate the potential drainage-reorganization mechanisms: antecedence, capture and overspill. Finally, we present a drainage-evolution model that documents key stages in the endorheic–exorheic transition, integrating the evidence for controls and mechanisms.

5.1. Controls on drainage evolution

Tectonic activity appears to have conditioned the landscape for drainage reorganization, with late Tortonian to Pliocene uplift concentrated around some of the margins (north, east and south) of the Douro Cenozoic Basin (Pyrenean and Iberian orogens, and Central System) (e.g., de Vicente et al., 2011, 2018; Cunha et al., 2019), causing progressive westward tilting and the formation of a regional topographic gradient (Fig. 1). This tectonically induced gradient would have primed the endorheic DCB for drainage re-organisation along its western margin, an area encompassing the present upper reaches of the Lower Douro. Differential uplift and the creation of regional topographic gradients is a commonly cited mechanism for enhancing base-level lowering and driving headward erosion, leading to capture-related reorganization (e.g. Stokes and Mather, 2003). In the case of the Douro, although uplift caused regional westward tilting, the relative base-level positions between the ancestral system routed to the Atlantic coast (Mondego Cenozoic Basin) and the endorheic DCB remained stable. As such, differential regional uplift would have been of insufficient magnitude to have driven or accelerated headward erosion by the ancestral Atlantic drainage system.

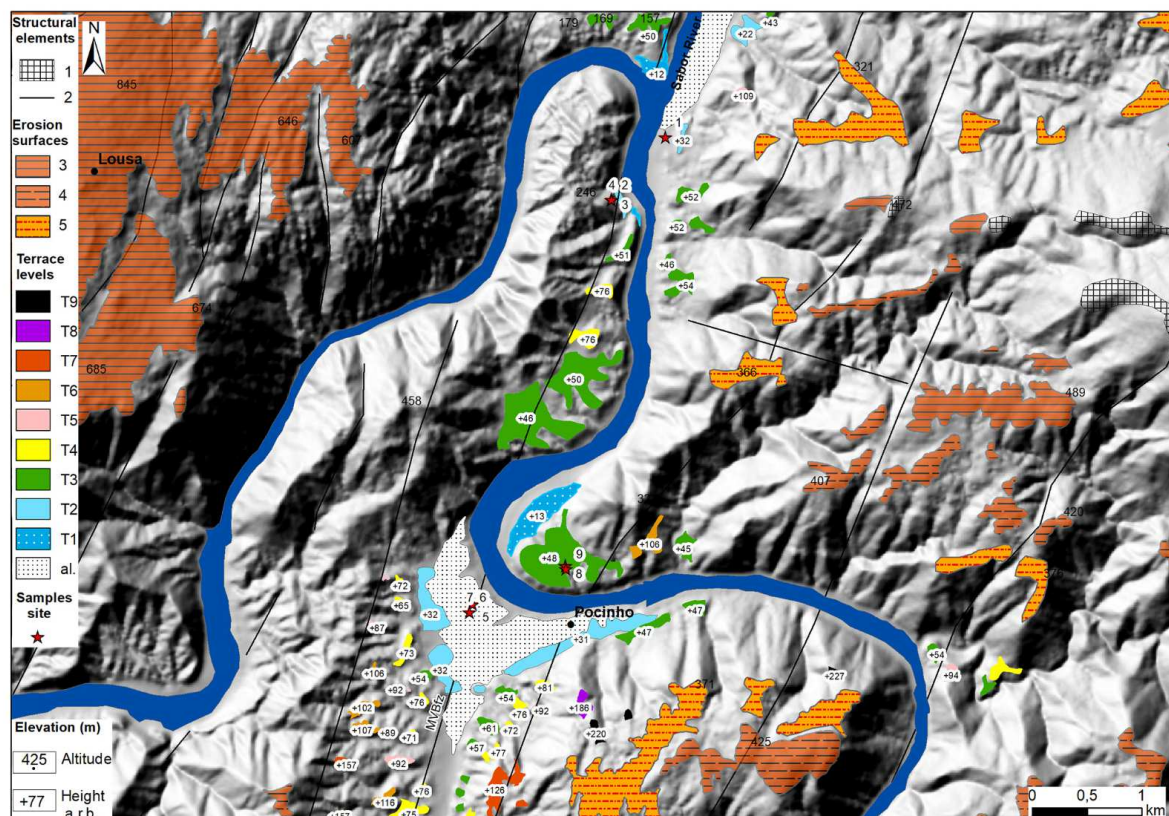


Fig. 10. Geomorphological map of the Pocinho terrace staircase. 1 - resistant reliefs on quartzites; 2 - tectonic lineament; 3 - North Iberian Meseta planation surface (NIMPS); 4 - degraded or displaced NIMPS; 5 - fluvial level at 365–371 m a.s.l. (+262 to +268 m a.r.b.); T9 to T1 - terrace levels; al. - alluvial plain. Sampled sites (numbered stars): 1 - Ribeira da Barca (sample 102217); 2, 3 and 4 - Quinta do Vale Meão (102211, 102212 and 102219); 5, 6 and 7 - Cortes da Veiga (102218, 102220 and 102221); 8 and 9 - Pocinho (102215 and 102216).

Climate during the period of drainage reorganization appears to have been hot and humid (early–mid Pliocene), becoming progressively cooler (late Pliocene – Pleistocene). The hot and humid conditions are especially important hydrological factors for the conditioning of a potential drainage reorganization. For example, increased water availability under hot and humid conditions would have resulted in expansion and rising levels of lakes within the endorheic DCB. When combined with the regional westward tectonic tilt, this would have made the western margin of the DCB a potential region for overflow/spill. Equally, the hot humid climate would have resulted in elevated flood discharges within the ancestral Atlantic drainage system, providing conditions conducive for headward erosion into, and potential capture of, the DCB.

Rock strength is variable throughout the Douro catchment, providing a factor that will either have enhanced or impeded fluvial incision. Higher rock strength is typically associated with the basement, and especially the granite, which dominates the WMDCB region where the reorganization is considered to have occurred. Studies by [Struth et al. \(2019\)](#) of knickpoint-migration rates from the upper reach of the Lower Douro through the WMDCB and up into the DCB during the drainage re-organisation period clearly demonstrate the incision – rock strength suppression and enhancement relationship. However, headward erosion by knickpoint migration requires base-level lowering. This could have been achieved by differential tectonic uplift between the Lower Douro and the DCB but, as previously stated, this was low–negligible, despite the regional westward tilting due to intraplate

compression. Eustatic base-level variability can also play a role in fluvial incision and knickpoint migration, for a marine-influenced system with a narrow offshore continental shelf. However, a sustained and large amplitude base level lowering would be required ([Bridgland and Westaway, 2012, 2014](#)). Global sea levels during the Pliocene were ~25–40 m above that at present and subsequently lowered in the Pleistocene as global climate cooled. Thus, eustatic base-level lowering might have some merit as a potential driver of headward erosion but again the occurrence of resistant granite in the upper Lower Douro and the WMDCB ([Fig. 1](#)) would have impeded migration, limiting the potential for capture-related reorganization. However, granite can be subject to significant amounts and rates of incision if flood discharges exploit favourable geological discontinuity patterns. [Antón et al. \(2015\)](#) recorded up to 100 m of fluvial incision and several hundreds of metres of headward erosion, when ‘normal’ winter floods were re-directed over a natural bedrock spillway during the construction of a dam on a Douro tributary within the WMDCB. This example is in keeping with an overspill process. Interestingly, cosmogenic exposure dating of fluvially eroded straths within bedrock river gorge reaches of the WMDCB by [Antón et al. \(2012\)](#) suggests exceptionally high incision rates within granite, again illustrating the potential for resistant bedrock to be rapidly eroded under favourable conditions.

5.2. Antecedence, capture or overspill mechanisms?

For antecedence to have occurred requires a pre-existing drainage

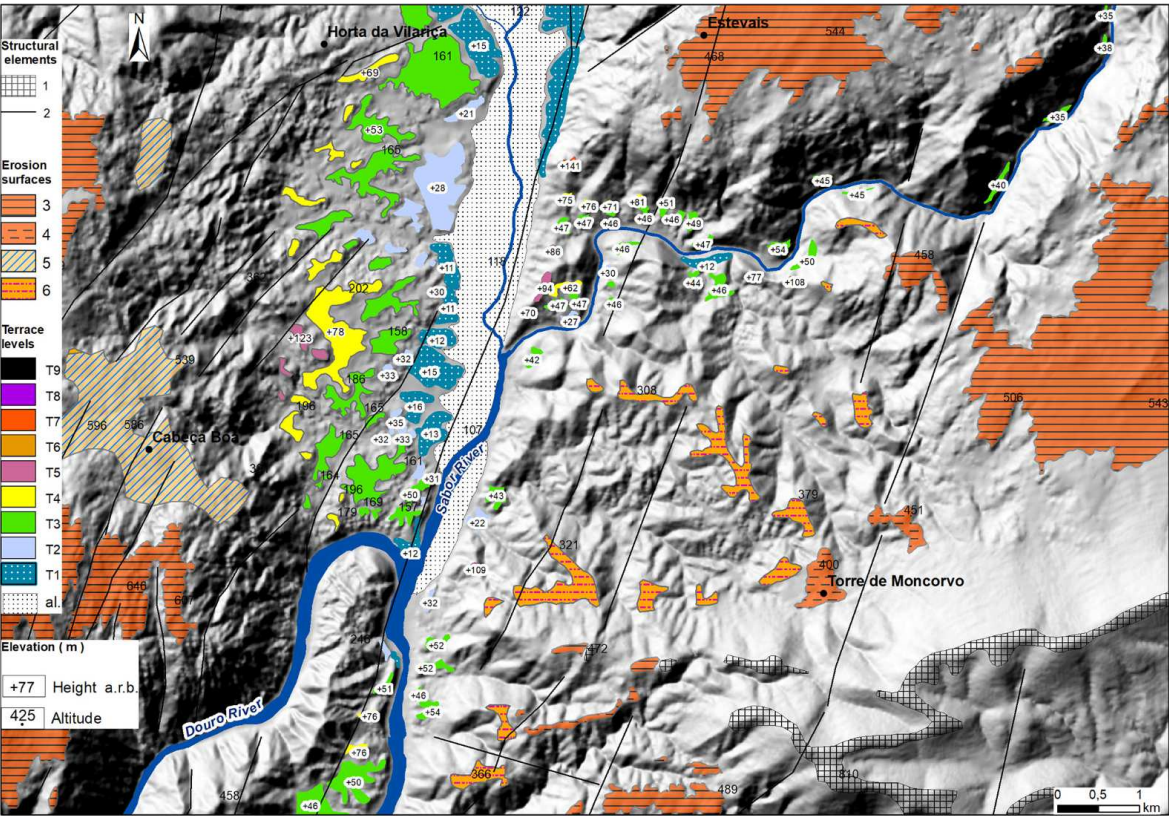


Fig. 11. Geomorphological map of the sector comprising the confluence of the Sabor with the Douro. 1 - residual quartzite upland; 2 - tectonic lineament; 3 - North Iberian Meseta planation surface (NIMPS), displaced by the Manteigas-Vilarica-Bragança fault zone; 4 - degraded remains of the NIMPS; 5 - fluvial level at 650–600 m a.s.l.; 6 - fluvial level at ~365 m a.s.l. (~+262 m a.r.b.); T9 to T1 - terrace levels; al. - alluvial plain.

network to have incised and kept pace with uplift (e.g. Stokes and Mather, 2003). This would require that an exorheic system traversing Iberia to the Atlantic coast to have already been established as part of the earliest mid-Neogene drainage pattern. Clearly this was not the case, since distinct endorheic (Douro Cenozoic Basin) and exorheic (Mondego Cenozoic Basin) systems were functioning as separate drainage systems in (and before) the early Pliocene. Tectonic uplift occurred in the study area during the generation of the Mountains and Plateaus of Northern Portugal, from the middle Tortonian (~9.6 Ma) to the mid-late Pliocene (~3.7 Ma). However, the sedimentary,

geomorphological and tectonic evidence suggests most of the regional drainage was received at that time by the endorheic Douro Cenozoic Basin (e.g., Pais et al., 2012; Cunha, 2019). Thus the endorheic to exorheic drainage reorganization would seem to be more readily explained in terms of 'bottom-up' capture-related headward fluvial erosion of an Atlantic draining river or by 'top-down' basin-overspill mechanisms.

River capture studies routinely document changes in sediment facies, palaeoflow evidence, provenance and grain-size characteristics (e.g. Harvey and Wells, 1987). For example, a system that has captured

Table 2
Main sedimentary characteristics of the studied Douro River terrace levels (Barca d'Alva to the Sabor River confluence). For all terraces, gravel composition is dominated by quartzite (milky quartz appearance), whereas sands are typically of arkosic composition.

| Geomorphic level | Elevation a.r.b (m) | Maximal thickness (m) | Sedimentary characteristics |
|------------------|---------------------|-----------------------|---|
| T3 terrace | +53 to +48 | ca. 12 | Boulder gravel terrace, > 10 m-thick at Barca d'Alva, clast-supported, with rare interbedded sand lenses (Fig. 12). At Qta. do Vale Meio there are fluvial channel fills of brown-reddish boulder gravels, which include lenses of coarse to medium sands and rare coarse silt levels. Boulders of quartzite up to 55 cm in diameter are common. Contains archaeological remains of Acheulian industries. |
| T2 terrace | +34 to +27 | ca. 18 | Clast-supported cobble gravels, with interbedded sand lenses. |
| T1 terrace | +17 to +13 | 11 | Comprises only compact fine to very fine sands and sandy silts of pale fawn colour. There is fossil root bioturbation at the top of channel filling sequences (1–2 m thick) and small pedogenic carbonate concretions within the topmost deposits. |
| Alluvial plain | +10 | 10 | Whitish fine friable sands (e.g. recent exposure made by channel down-cutting just downstream of the Pocinho dam). |
| Modern river bed | +0 | ca. 2 | Coarse sands to gravels, overlying the Variscan basement. |

Table 3

Geographic coordinates, burial depth, water content and radionuclide activities used for dose rate calculation.

| NLL and field codes | Geographic coordinates (M, P) | Sample altitude (m) | Burial depth (cm) | Water content (%) | U-238 (Bq kg ⁻¹) | Ra-226 (Bq kg ⁻¹) | Th-232 (Bq kg ⁻¹) | K-40 (Bq kg ⁻¹) |
|---------------------|-------------------------------|---------------------|-------------------|-------------------|------------------------------|-------------------------------|-------------------------------|-----------------------------|
| 102211 T1VM-1 | N41°10'0.65" W7° 6'53.46" | 119.5 (+13.5) | 655 | 15 (3;30) | 57 ± 8 | 47.7 ± 0.8 | 59.3 ± 0.9 | 644 ± 13 |
| 102212 T1VM-6 | N41°10'0.65" W7° 6'53.46" | 123 (+17) | 285 | 15 (4;24) | 41 ± 6 | 41.1 ± 0.6 | 47.8 ± 0.7 | 567 ± 11 |
| 102213 BALVA-2 | N41°01'53" W006°55'50" | 158 (+45 m) | 300 | 7 (1;11) | 18 ± 6 | 17.1 ± 0.5 | 19.7 ± 0.5 | 826 ± 16 |
| 102214 BALVA-1 | N41°01'24.09" W6°56'9.87" | 178 (+65 m) | 40 | 10 (5;20) | 43 ± 6 | 49.3 ± 0.7 | 66.4 ± 0.8 | 645 ± 12 |
| 102215 POCINHO-2 | N41°8'12.41" W7°7'13.12" | 151 (+45) | 800 | 12 (8;22) | 29 ± 7 | 15.5 ± 0.5 | 25.3 ± 0.6 | 1040 ± 20 |
| 102216 POCINHO-1 | N41°8'11.47" W7°7'13.01" | 153 (+47) | 600 | 8 (2;14) | 15 ± 5 | 13.8 ± 0.4 | 24.3 ± 0.5 | 755 ± 14 |
| 102217 RBARCA-1 | N41°10'19" W007°6'32" | 131 (+25) | 80 | 14 (13;26) | 38 ± 8 | 36.1 ± 0.7 | 55.3 ± 0.9 | 548 ± 12 |
| 102218 CVEIGA-1 | N41°07'59" W007°7'51" | 124 (+18) | 210 | 10 (6;23) | 31 ± 7 | 39.3 ± 0.7 | 50.1 ± 0.8 | 562 ± 12 |
| 102219 T1VM-7 | N41°10'0.65" W7° 6'53.46" | 125 (+19) | 90 | 17 (9;27) | 52 ± 8 | 51.2 ± 0.8 | 54.0 ± 0.9 | 592 ± 13 |
| 102220 CVEIGA-4 | N41° 8'1.40" W7° 7'49.80" | 123 (+17) | 200 | 8 (1;16) | 15 ± 6 | 22.5 ± 0.5 | 25.8 ± 0.6 | 756 ± 14 |
| 102221 POCINHO-0 | N41°07'59" W007°7'51" | 106 (+0) | 10 | 18 (17;21) | 145 ± 7 | 138.1 ± 1.3 | 67.6 ± 0.8 | 1273 ± 21 |
| 112233 BALVA-3 | N41°01'34" W006°56'31" | 160.5 (+47.5) | 1100 | 6 (10;13) | 40 ± 5 | 24.2 ± 0.5 | 39.9 ± 0.7 | 372 ± 9 |
| 112234 ALMENDRA-1 | N41°02'08" W007°0'10" | 113 (+0) | 5 | 25 (25;25) | 17 ± 5 | 24.8 ± 0.4 | 35.3 ± 0.6 | 639 ± 11 |

other drainage would show evidence of an enlarged catchment area, such as elevated sediment/water discharges being delivered downstream as the captured headwaters undergo upstream propagation of a fluvial incision wave as an adjustment to connection with a new and lower base level. The new larger system could also have lower downstream gradient, this being proportional to catchment size. The elevated sediment/water fluxes result in marked increases in grain size and changes to fluvial sedimentation styles (e.g. fluvial pattern change: meandering to braided). The addition of new areas to the catchment might also result in different rocks becoming available to headwaters and thus being recorded in downstream fluvial sediment deposited post-capture. In the Lower Douro, the early drainage is recorded as basin sedimentary infill adjacent to the coast (the Mondego Cenozoic Basin). Sedimentary and palaeogeographic reconstructions (Cunha, 2019) reveal a fine-grained sandy braided drainage system connected to the Atlantic with sediment provenance from localized coastal relief. This older drainage system thus lacks sedimentary evidence for capture.

The younger drainage system of the Lower Douro is recorded within inset sequences of river terraces. These terraces appear to have a temporal and spatial relationship to similar sequences upstream within the DCB. Sediment in these terraces sequences can be shown to have been transported downstream from the DCB into the Lower Douro towards the Atlantic coast (Silva et al., 2017). The terraces have similarities in sedimentary style and clast provenance that lack any unequivocal signature of capture (e.g., Harvey and Wells, 1987; Harvey et al., 1995). The continuity of terraces throughout the Douro catchment and the evidence for palaeoflow towards the Atlantic suggests that the drainage reorganization pre-dated terrace formation and correlate to the initiation of drainage to the coastal basin. Neither of these Lower Douro systems provides compelling evidence for capture-related drainage reorganization. If capture occurred, evidence for it is missing from the Lower Douro region, perhaps due to headward erosion by migrating knickpoints (e.g. Struth et al., 2019). If a capture-related wave of incision had propagated upstream along the WMDCB, its progress would have been limited by the resistant granite (e.g. Struth et al., 2019). However, the shape of the modern Douro long profile (Fig. 1) lacks any clear capture signature, instead having a morphology that reflects a combination of rock strength and adjustment after the exorheic drainage re-organisation occurred.

Overspill requires water to cross a catchment divide and is a widely cited explanation for drainage reorganization from many studies from tectonically elevated regions, with either mountainous relief and/or high-elevation plateaus (e.g. Grand Canyon: Douglass et al., 2009; Tibetan Plateau: Craddock et al., 2010). The process typically involves the growth of a lake, with an increase in water level which ultimately reaches the level of the catchment divide, over which the rising water spills. Classic examples of overspill come from the internally drained

Great Basin region of the SW USA, where lake-level changes linked to Pleistocene glacial/pluvial periods can be tracked by the positions of shoreline sequences (Tchakerian and Lancaster, 2002), with relationships to eroded canyons (gorges) and fluvial terraces downstream of overspill sites (e.g. Menges, 2008). In the present study, the DCB has a lacustrine configuration, but this evidence is derived from the sedimentary basin infill alone; geomorphological lake-shoreline terraces are absent. The DCB lake may have extended downstream through the WMDCB and into the study area (uppermost reach of the Lower Douro). However, evidence is restricted to fragmentary remnants of high-level (~665–365 m a.s.l.) surfaces of wide landscape extent (Figs. 4 and 5) but lacking sediments. The 650–600 m a.s.l. surface ends along the margins of the Montemuro-Marão upland in the Mesão Frio area (Fig. 4), with no downstream continuation, suggesting that this location was the limit of potential lacustrine extent, with the position of the drainage divide lithologically controlled by resistant granite and quartzite basement lithologies (Figs. 1, 2 and 4). Upstream of Mesão Frio, as far as Barca d'Alva, the dominance of weak metamorphic lithologies would have helped any overflow from a lake/ponded system to incise the erosional surface. Regional westward tectonic tilting (Section 5.1) may have helped promote ponding in this area. Furthermore, the warm and humid climate of the middle Pliocene would have promoted expansion and water-level rise in lake-systems, but improved age control is needed to establish this relationship with greater security. The inset ~575–365 m a.s.l. surface extends downstream beyond the Montemuro-Marão upland into the Lower Douro (Fig. 4). Its genetic origin is unclear but its inset nature and continuation downstream, where it is cut into granite bedrock (Figs. 1, 2 and 4), suggests that a fluvial origin is more likely. When the high-level surfaces are considered together, it seems likely that drainage reorganization occurred with the formation of the 650–600 m level; the formation of the 575–365 m level can be related to clear exorheic drainage. Collectively, evidence from the surface geomorphology, rock strength, tectonic and climatic relationships of these surfaces appear to integrate well and provide marginally more compelling evidence for overspill than for capture as the mechanism for drainage reorganization.

5.3. Endorheic to exorheic drainage transition in the Douro

Integration of the previously discussed controls and mechanisms for drainage evolution has allowed a conceptual model for the endorheic to exorheic transition of the Douro to be established. We consider this transition to have occurred in four stages:

Stage 1: The initial stage involved the development of the endorheic DCB. The DCB was formed and filled as an intermontane foreland basin. Along its western margin, it tapered into a high-elevation plateau formed from regionally eroded igneous and metamorphic basement,

Table 4
Summary of the luminescence ages obtained from the Douro River samples of the study area. The age considered more reliable (Qz versus Kf, used as the dosimeter) is in Bold.

| NLL and field codes | Site | Sampled level | Sedimentary facies | Sample altitude (m) | Burial depth (cm) | Equivalent dose (Gy) | Dose rate (Gy/ka) | n | OSL age (ka) |
|---------------------|-------------------------------|---------------------|--|---------------------|-------------------|--------------------------------------|--------------------------------------|-------|---|
| 102211 TIVM-1 | Quinta do Vale Meio (Pocinho) | T1 (left bank) | Sandy silt with carbonate concretions | 119.5 (+13.5) | 655 | 88 ± 2 (Qz) | 3.48 ± 0.15 (Qz) | 28 | 25.2 ± 1.3 (Qz) |
| 102212 TIVM-6 | Quinta do Vale Meio (Pocinho) | T1 (left bank) | Sandy silt with carbonate concretions | 123 (+17) | 285 | 51 ± 2 (Qz) | 3.02 ± 0.13 (Qz) | 32 | 17.0 ± 1.0 (Qz) |
| 102213 BALVA-2 | Barca d'Alva | T3 (right bank) | Coarse sand | 158 (+45 m) | 300 | > 200 (Qz) > 1000 (Kf) | 3.18 ± 0.14 (Qz) 4.03 ± 0.15 (Kf) | 33 | > 60 (Qz) > 270 (Kf) |
| 102214 BALVA-1 | Barca d'Alva | T3 top (left bank) | Brown coarse silt (aeolian cover) | 178 (+65 m) | 40 | 122 ± 6 (Qz) 129 ± 6 (Kf) | 3.86 ± 0.17 (Qz) 4.64 ± 0.17 (Kf) | 45 4 | 31 ± 2 (Qz) 28 ± 2 (Kf) |
| 102215 POCINHO-2 | Pocinho | T3 (right bank) | Coarse sand | 151 (+45) | 800 | > 200 (Qz) > 1000 (Kf) | 3.66 ± 0.16 (Qz) 4.57 ± 0.17 (Kf) | 37 | > 55 (Qz) > 230 (Kf) |
| 102216 POCINHO-1 | Pocinho | T3 (right bank) | Medium sand | 153 (+47) | 600 | > 200 (Qz) > 1000 (Kf) | 3.02 ± 0.14 (Qz) 3.91 ± 0.15 (Kf) | 33 | > 70 (Qz) > 290 (Kf) |
| 102217 RBARCA-1 | Ribeira da Barca (Pocinho) | T2 (right bank) | Very coarse sand | 131 (+25) | 80 | 176 ± 8 (Qz) 382 ± 21 (Kf) | 3.10 ± 0.13 (Qz) 3.90 ± 0.14 (Kf) | 24 7 | 57 ± 4 (Qz) 98 ± 7 (Kf) |
| 102218 CVEIGA-1 | Cortes da Veiga (Pocinho) | T1 (left bank) | Fine sand with some carbonate concretions | 124 (+18) | 210 | 40.6 ± 1.4 (Qz) | 3.15 ± 0.14 (Qz) | 23 | 12.9 ± 0.8 (Qz) |
| 102219 TIVM-7 | Qia Vale Meio (Pocinho) | T1 (left bank) | Very fine sand with bioturbation and carbonate concretions | 125 (+19) | 90 | 55 ± 2 (Qz) | 3.30 ± 0.14 (Qz) | 18 | 16.7 ± 1.0 (Qz) |
| 102220 CVEIGA-4 | Cortes da Veiga (Pocinho) | T1 (left bank) | Fine sand | 123 (+17) | 200 | 46.0 ± 2.1 (Qz) | 3.11 ± 0.14 (Qz) | 30 | 14.8 ± 1.0 (Qz) |
| 102221 POCINHO-0 | Cortes da Veiga (Pocinho) | modern river bed | Very coarse sand | 106 (+0) | 10 | 1.36 ± 0.37 (Qz) | 6.47 ± 0.27 (Qz) | 19 | 0.21 ± 0.06 (Qz) |
| 112233 BALVA-3 | Barca d'Alva | T3 base (left bank) | Fine sand lense in gravels | 160.5 (+47.5) | 1100 | > 1000 (Kf) | 2.80 ± 0.10 (Kf) | 6 | > 360 (Kf) |
| 112234 ALMENDRA-1 | Barca d'Alva | modern river bed | Fine sand | 113 (+0) | 5 | 0.020 ± 0.006 (Qz) 10.0 ± 1.7(Kf) | 2.66 ± 0.09 (Qz) 3.51 ± 0.11 | 30 12 | 0.020 ± 0.006 (Qz) 2.80 ± 0.50 (Kf) |



Fig. 12. Photo illustrating the T3 terrace deposits at Barca d'Alva, Douro right bank. The large lense of medium sand located in the middle part of the exposure was sampled for OSL dating (sample 102213; BALVA-2; 158 m a.s.l., +45 m a.r.b.).

shaped between the early Cretaceous and the late Cenozoic. The MPNP and NIMPS mark the western extent of the endorheic DCB. Evidence for a contemporaneous Atlantic drainage system is lacking.

Stage 2: This marks the onset of drainage reorganization. Westward tectonic tilting of the DCB and elevated Pliocene hydrological conditions resulted in ponding and the development of an extended lake system, resulting in a high-level fluvial surface at 600–650 m a.s.l. This surface, although now fragmentary, seems to have had a widespread spatial extent within the tectonic depressions along the westernmost margin of the endorheic system (i.e. between Mesão Frio and Barca d'Alva in the uppermost reach of the Lower Douro). The endorheic system was confined by a drainage divide in resistant geology at Mesão Frio. The overspilling of this divide as a result of elevated water levels initiated erosion and breaching of this barrier, marking the onset of exorheic drainage (Fig. 14).

Stage 3: Headward erosion, initiated by the overspill and subsequently progressing upstream into the depression and further into the WMDCB was moderated by the variable resistance to erosion of the metamorphic and granite basement (Fig. 15). The 575–365 m a.s.l. surface within the uppermost Lower Douro depression marks the establishment of exorheic drainage, but a full connection with the Mondego Cenozoic Basin was probably yet to be established. Age modelling using the incision rates of dated river terrace deposits (Stage 4) suggests that the 575–365 m a.s.l. surface was developed during the early Pleistocene (~1.8 Ma). This implies that drainage reorganization occurred during the Pliocene.

Stage 4: This final stage of drainage evolution was marked by enhanced vertical downcutting, during which river terrace staircases developed, both within the study area (Figs. 6, 9, 10, 11) and upstream in the DCB (Silva et al., 2017). This stage, with rapid incision, represents

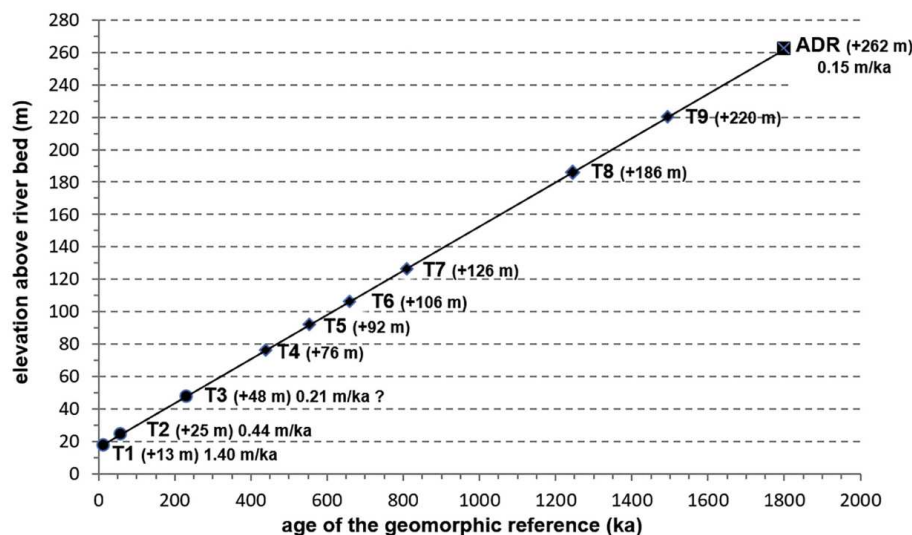
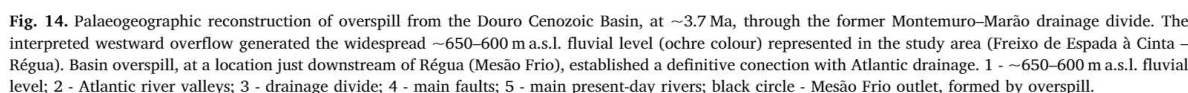


Fig. 13. For the Pocinho terrace staircase, age estimates for the uppermost deposits of T1, T2 and T3 are plotted against height a.r.b. (black circles). Extrapolation suggests probable ages for the older fluvial straths (black lozenges) and for the ancestral Douro valley-floor (ADR; black square). Thus, a probable age of 1.8 Ma is likely for the onset of enhanced fluvial incision.



Basin towards the west, and (ii) a major climate change by ~ 3.7 Ma (transition to a wetter climate stage), from the generally dry and hot climate during the Miocene and Zanclean to the humid and hot climate of the late Zanclean to Piacenzian.

The latest stage (last ~ 2 Ma), marked by enhanced fluvial incision, is related to continuous regional low crustal uplift, cooler climatic minima and associated lowering of sea level. The progressive extension of the Atlantic coastal drainage system into the Douro Cenozoic Basin drove a progressive increase in discharge, and thus incision, in the Lower Douro, explaining the progressive entrenchment of the valley. Bedrock (strath) terraces can be correlated along the longitudinal profile of the Lower Douro over considerable distances, pointing to almost synchronous development rather than the result of a unique wave of incision transmitting upstream during the Pleistocene.

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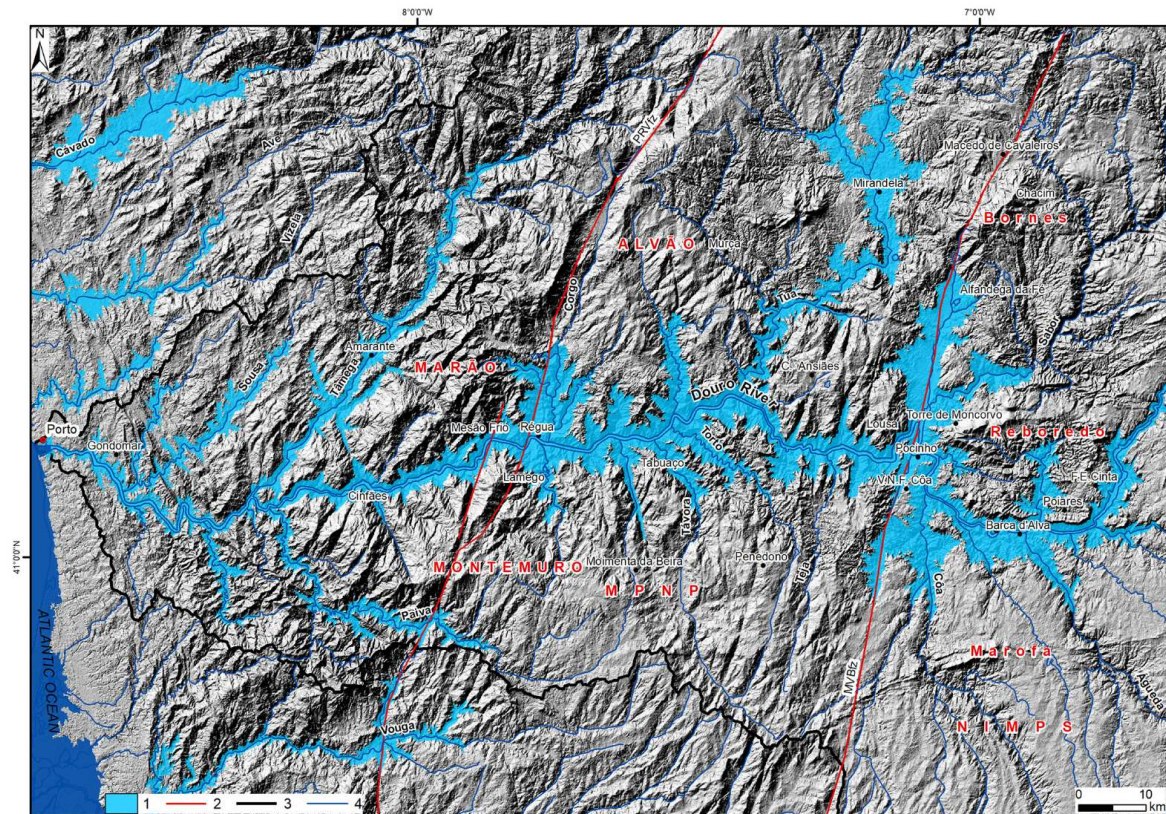


Fig. 15. Palaeogeographic reconstruction of the ancestral Douro River and its main tributaries, at ~3.6–2 Ma (just before the rapid fluvial incision began – Stage 4). 1 – 575–365 m a.s.l. fluvial level; 2 – main faults; 3 – drainage divide; 4 – main present-day rivers.

Declaration of Competing Interests

The authors declare no conflict of interest. The funders had no role in the design of the study; in the collection, analyses, or interpretation of data; in the writing of the manuscript, or in the decision to publish the results.

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Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.gloplacha.2019.102985>.

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