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Late Variscan deformation in the Iberian Peninsula; a late feature in the Laurentia–Gondwana dextral collision

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7 Abstract The Late Variscan deformation event in Iberia is characterized by an intraplate deformation regime 8 induced by the dextral oblique collision between Laurentia 9 and Gondwana. This episode in Iberia is characterized by 10 NNE-SSW brittle to brittle-ductile strike-slip faults, which 11 are considered by the classic works as sinistral strike-12 slip faults. However, the absence of Mesozoic formations 13 constraining the age of this sinistral kinematics led some 14 authors to consider it as the result of Alpine reworking. 15 Structural studies in Almograve and Ponta Ruiva sectors 16 (SW Portugal) show that NNE-SSW faults have a sinistral 17 kinematics and are occasionally associated with E-W dex-18 19 tral shears. Moreover, this kinematics is related to the late deformation episodes of Variscan orogeny. In Almograve 20 sector, the Late Variscan structures are characterized by 21 NNE-SSW sinistral kink bands, spatially associated with 22 E-W dextral faults. These structures are contemporaneous 23 and affect the previously deformed Carboniferous units. 24 The Ponta Ruiva Sector constrains the age of deformation 25

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because the E-W dextral shears affect the Late Carbonif-26 erous (late Moscovian) units, but not the overlying Trias-27 sic series. The new data show that the NNE-SSW and the 28 E-W faults are dynamically associated and result from the 29 Late Variscan deformation episode. The NNE-SSW sin-30 istral faults could be considered as second-order domino 31 structures related to first-order E-W dextral shears, linked 32 to Laurentia-Gondwana collision during Late Carbonifer-33 ous-Permian times. 34

Keywords Iberia Variscan orogeny · Late Variscan deformation · South Portuguese Zone · Kink bands

Introduction

A complex network of major shear zones was developed 38 during the last stages of intracontinental deformation of 39 the Variscan orogeny. This Late Variscan deformation epi-40 sode was considered the result of internal deformation 41 along first-order E-W dextral shear zones (Arthaud and 42 Matte 1975, 1977). Such kinematics is often considered a 43 pervasive feature of most of the Variscan orogenic evolu-44 tion (Ribeiro et al. 1995; Shelley and Bossière 2000, 2002; 45 Ribeiro 2002; Ribeiro et al. 2007; Martínez Catalán 2011; 46 Nance et al. 2012; Dias et al. 2016). In the Iberian Mas-47 sif, this Late Variscan deformation gave rise to some of the 48 most important observed basement faults (Ribeiro 1974; 49 Iglesias and Ribeiro 1981), like the NNE-SSW Vilariça 50 and Régua-Chaves-Verin faults in NW Iberia (Ribeiro et al. 51 1990; Marques et al. 2002; Moreira et al. 2010; Dias et al. 52 2013). Although several works focus on this major event, 53 there are still some doubts concerning the kinematics and 54 the timing of the deformation. Such controversy mainly 55 results from the strong reworking of the Late Variscan 56



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